

CO₂-SO₂ clathrate hydrate formation on early Mars

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Abstract. Most sulfate minerals discovered on Mars are dated no earlier than the Hesperian. We showed, using a 1-D radiative-convective-photochemical model, that clathrate formation during the Noachian would have buffered the atmospheric CO₂ pressure of early Mars at ~2 bar and maintained a global average surface temperature ~230 K. Because clathrates trap SO₂ more favorably than CO₂, all volcanically outgassed sulfur would have been trapped in Noachian Mars cryosphere, preventing a significant formation of sulfate minerals during the Noachian and inhibiting carbonates from forming at the surface in acidic water resulting from the local melting of the SO₂-rich cryosphere. The massive formation of sulfate minerals at the surface of Mars during the Hesperian could be the consequence of a drop of the CO₂ pressure below a 2-bar threshold value at the late Noachian-Hesperian transition, which would have released sulfur gases into the atmosphere from both the Noachian sulfur-rich cryosphere and still active Tharsis volcanism. Our hypothesis could allow to explain the formation of chaotic terrains and outflow channels, and the occurrence of episodic warm episodes facilitated by the release of SO₂ to the atmosphere. These episodes could explain the formation of valley networks and the degradation of impact craters, but remain to be confirmed by further modeling.

1 Thermodynamic modelling of CO₂-SO₂ clathrates in Martian atmospheric conditions

Using an existing thermodynamical approach, a model has been used to extrapolate to low temperatures existing data on the composition of CO₂-SO₂ clathrates. At 200-220 K, the enrichment factor of SO₂ in the clathrate with respect to gas is found to be very high, at least ≈100.

2 General scenario of Mars early evolution

Based on the previous result, we have shown (Chassefière et al., 2013) that the formation of CO₂-SO₂ clathrates at Noachian and Hesperian times could have played an important role in controlling and stabilizing the level of volcanic sulfur in the atmosphere, as well as the level of atmospheric CO₂ for earliest times. If the CO₂ pressure exceeded a threshold of 2 bar due to an efficient Noachian

volcanism, the CO₂ in excess of 2 bar could have been stored in the cryosphere under the form of CO₂ clathrates. Indeed, due to the increasing albedo of the atmosphere through Rayleigh scattering for increasing CO₂ pressure above 1 bar, the surface temperature induced by a p_{CO₂}>~2 bar CO₂ atmosphere is smaller than the equilibrium temperature of clathrates, resulting in a saturation of CO₂ and its condensation under the form of clathrates. Such a clathrate buffer would have maintained the CO₂ pressure close to ~2 bar, and the surface temperature close to 230 K, resulting in a cold Mars at the Noachian. The cryosphere would have trapped all the sulfur released by volcanism under the form of sulfur-rich (enriched by a factor ~100-500 with respect to the gas phase) CO₂-SO₂ clathrates.

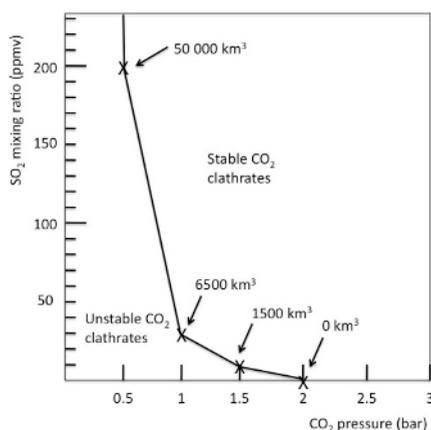


Fig. 1. Equilibrium SO₂ mixing ratio as a function of CO₂ pressure. The lava volumes required to give rise to the corresponding SO₂ mixing ratio are indicated on the figure.

For a surface pressure smaller than 2 bar and progressively decreasing, atmospheric sulfur is no longer trapped in clathrates (Fig. 1), and an increasing fraction of the atmospheric SO₂, up to 10 ppm at p_{CO₂}=1.5 bar, and 30 ppm at p_{CO₂}=1 bar, can remain in the atmosphere. If the amount of atmospheric SO₂ exceeds these thresholds, the cooling effect of sulfate aerosols results in a decrease of surface temperature below the clathrate equilibrium temperature and the condensation of the SO₂ in excess in CO₂-SO₂ clathrates. During this period, significant amount of SO₂ was converted into sulfate aerosols, which further settled down to the surface and possibly led to the formation of sulfate minerals. This sulfur may have been released, not only by volcanoes but also (and more continuously) by the SO₂-rich cryosphere formed when p_{CO₂}>~2 bar. Through this mechanism, the cryosphere may have lost at this stage some of the volcanic SO₂ stored at earlier times, released back to the atmosphere. In this time range (2 to 1 bar p_{CO₂} range), the surface temperature has been buffered, during and after episodes of sulfur release, at a temperature from 230 K (p_{CO₂}=2 bar) to 210 K (p_{CO₂}=1 bar). For p_{CO₂}<1 bar, all the released SO₂ remains in the atmosphere, with no more trapping in clathrates.

Accordingly, we have proposed the hypothesis that the formation of sulfate minerals, which have been observed from orbit and are formed during the Hesperian (Bibring et al., 2006), could have been triggered by a fall of early Mars atmosphere pressure below 2 bar and that this change occurred close to the late Noachian/Hesperian transition.

3 Major geomorphological processes compatible with this scenario

Two majors geomorphological units are compatible with our scenario:

1) The presence of a thick Mars-wide sedimentary formation of sulfate, including in very uncommon places at high topography, such the Interior Layer Deposit (ILD) at the top of Valles Marineris. Eolian deposition was proposed from geomorphological arguments. For ILD in Valles

Marineris, the formation as thick as 5 km must be done in 400 Myrs, after the tectonic opening at 3.9 Ga and the formation of the floor 3.5 Ga. In our scenario, 20 m to 50 m GEL sulfur particle could have precipitated directly from atmosphere in 40 to 1000 Myrs, in agreement with the observation.

2) The chaotic terrains, at equatorial region, could have been formed by disruption of the CO₂ clathrate in the past at the late Hesperian/Amazonian period. This interpretation is compatible with our scenario since the formation of clathrate must have been global during the Noachian, it should also affect equatorial region. The reason for such a disruption has been debated, including climatic change, internal heat flux, fracture propagation, and seismic activity. We propose here that the disruption may be due to the pressure decrease below 2 bar. In our scenario, a GEL of CO₂-SO₂ clathrate ranging from 290 m to 720 m depth has been destabilized. Scaling this volume to the province of the chaos (1/10 surface of Mars) is compatible with the typical height loss in chaos ~3000 m (Fig. 2)

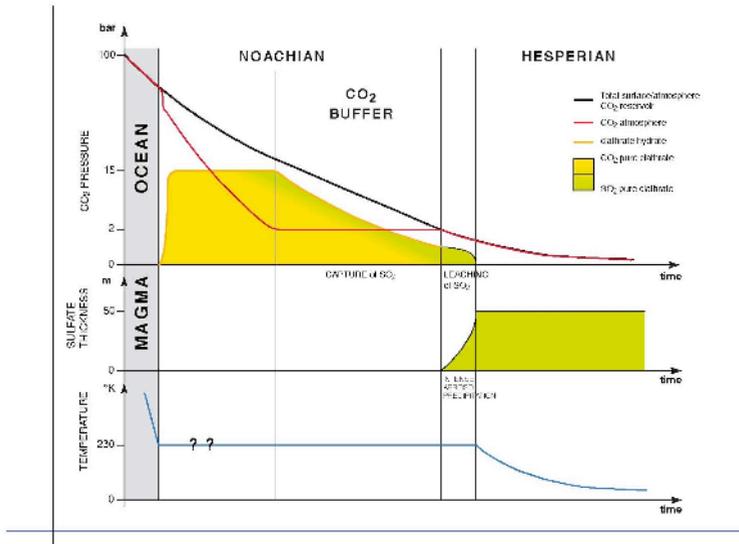


Fig. 2. Schematic representation of sulfur and CO₂ reservoir during the Noachian and Hesperian period. The total volatile surface/atmosphere CO₂ reservoir is decreasing due to thermal escape and possible formation of carbonates in the subsurface, from an initial budget ~100 bar. After magma ocean crystallization, CO₂ and water condense in form of clathrate (15 bar of CO₂ at maximum can be stored in form of clathrate when reacting with the total Martian water content of 1000 m GEL). After volcanic release of S, atmospheric SO₂ is systematically enriched in the condensating clathrate, triggered either by climatic changes or by global cooling due to aerosol formation. The CO₂ clathrate acts as an atmospheric buffer when the atmospheric pressure reaches 2 bar. After the complete consumption of CO₂ clathrate, the atmospheric pressure can drop below 2 bar and massive release of SO₂ from the clathrate induces the condensation of aerosols and precipitation as a Mars-wide sulfate layer (50 m represent the total SO₂ outgassed from Tharsis formation). This scenario represents the equilibrium state of surface/atmosphere but significant departure may have happened due to large impacts.

References

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